



湖泊微生物反硝化过程及速率研究进展

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摘要: 湖泊中微生物介导的反硝化过程对于区域乃至全球的气候环境变化有着深远的影响。因此, 研究湖泊微生物反硝化过程及速率有助于我们深刻理解湖泊氮元素生物地球化学循环规律, 全面认识湖泊生境对全球氮循环的贡献。本文综述了湖泊生境中反硝化过程(包括典型的反硝化过程及与其他物质循环耦合的反硝化过程, 如与有机氮耦合的共反硝化作用、与碳循环耦合的硝酸盐/亚硝酸盐依赖型厌氧甲烷氧化、与铁循环耦合的硝酸盐依赖型铁氧化、与硫循环耦合的硝酸盐还原硫氧化)的速率、驱动微生物及其影响因素。最后对湖泊反硝化过程研究现状和未来发展方向提出总结与展望。

关键词: 湖泊, 微生物, 反硝化, 速率, 影响因素

内陆水生系统在全球氮循环中扮演着重要的角色。其面积不足陆地面积的 4%, 却贡献了约 50% 的陆地氮流失^[1]。其每年运移、矿化、埋藏的氮总量(0.1 Pg, Pg=10¹⁵ g)约等于人类工业活动导致的陆地氮输入(0.1 Pg), 接近海洋总固氮量(0.14 Pg)^[2](表 1)。因此, 研究内陆水生系统的氮循环过程能够帮助我们科学、全面地认识全球氮循环。

湖泊作为内陆水体的主要组成部分, 约占其总面积的 83%, 且湖泊单位面积的氮去除速率与河流相当, 远高于其他内陆水体(如地下水), 在内陆水体氮循环中扮演着举足轻重的角色^[1]。微

生物作为湖泊氮循环的主要驱动者, 参与推动湖泊氮固定、运移、循环转化以及埋藏等多个重要的氮循环过程。因此, 研究湖泊微生物驱动的氮循环过程有助于我们深刻理解湖泊氮循环的机制, 全面认识湖泊氮循环在全球生物化学循环中的作用。

由于承受着陆源过量的营养盐输入与持续增长的 NO₃⁻型氮沉降, 湖泊生境生态功能有着明显的退化趋势^[13-14]。反硝化过程在减轻湖泊氮载荷方面起着不可替代的作用, 能够将湖泊中的硝态氮转化为 N₂O 或 N₂。在某些湖泊中, 氮的去除率高达 36% (与氮输入总量相比)^[15]。然而, N₂O

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是一种强大温室气体, 其增温潜势是 CO₂ 的约 300 倍, 还会在平流层进一步氧化为 NO 破坏臭氧层, 因此, N₂O 排放速率也越来越受到人们的

广泛关注^[16-17]。除了典型的反硝化过程, 近年来与有机氮、甲烷、铁、硫等元素或化合物耦合的反硝化新途径也开始进入学者们的视野(图 1), 这

表 1. 全球氮通量
Table 1. Global nitrogen fluxes

Pathways	Nitrogen fluxes (Tg N/yr)	Ranges	References
Sources	483		
Biological nitrogen fixation of inland waters	74	\	[3]
Nitrogen deposition of inland waters	23	\	[4]
Terrestrial nitrogen input of inland waters	7	\	[4]
Terrestrial nitrogen fixation	84	40–127	[5–6]
Marine nitrogen fixation	134	94–175	[7–8]
Lighting	5	\	[5]
Industrial nitrogen fixation	105	100–110	[5, 9]
NO _x from combustion	38	25–52	[5]
Agricultural nitrogen fixation	43	32–50	[5, 10]
Sinks	521		
Inland waters denitrification	110	39–216	[1]
Inland waters N ₂ O emission	0.22	0.15–0.28	[11]
Terrestrial denitrification	130	58–175	[1]
Marine denitrification	245	107–331	[5, 7]
Terrestrial N ₂ O emission	12	\	[2]
Marine N ₂ O emission	4	\	[2]
Marine nitrogen burial	20	16–25	[2, 5, 12]

“\” means no reference data.

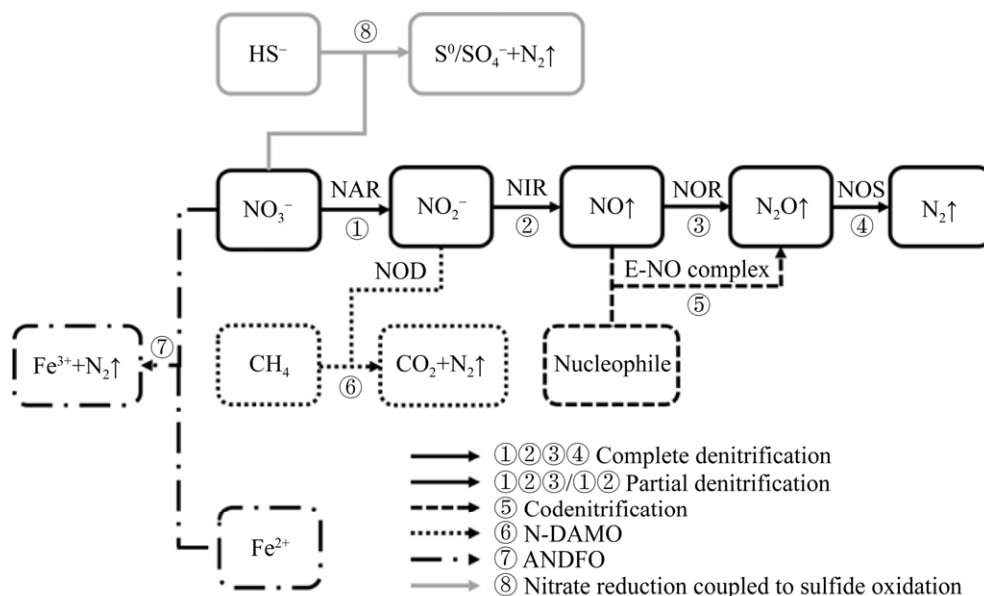


图 1. 反硝化与其他元素循环耦合模式图

Figure 1. Typical denitrification processes and those coupled with other elemental cycles. N-DAMO: Nitrite-Dependent Anaerobic Methane-Oxidization; ANDFO: Anaerobic Nitrate-Dependent Fe(II) Oxidation; NAR: nitrate reductase; NIR: nitrite reductase; NOR: nitric oxide reductase; NOS: nitric oxide synthase; E-NO complex: enzyme (E) bound NO complexes; NOD: nitric oxide dismutase.

些新途径对于湖泊的氮去除过程也有一定程度的贡献^[18-21]。本文将综述湖泊生境中不同类型的反硝化过程,介绍其速率及驱动微生物的群落组成,并综合微生物代谢机理分析反硝化过程的影响因素,总结现有研究的不足,展望未来研究的方向。

1 湖泊中的反硝化过程

1.1 硝酸盐还原为亚硝酸盐

反硝化作用中的硝酸盐还原指的是微生物通过呼吸作用将硝酸盐转化为亚硝酸盐的异化还原过程。硝酸盐的异化还原过程主要依靠兼性厌氧的细菌、古菌和真核生物(如 α 、 β 和 γ 变形菌纲的细菌和一些嗜盐古菌)等完成^[22-23]。相对于其他氮循环过程,硝酸盐还原是湖泊中亚硝酸盐的主要贡献者。而亚硝酸盐作为中间产物继而通过三种途径被消耗^[22,24-25]:(1)通过反硝化作用被还原为气态产物(N_2O 或 N_2);(2)通过硝酸盐异化作用被还原成铵(DNRA,又称硝酸盐铵化);(3)通过厌氧氨氧化(anammox)过程被还原为氮气。

1.2 亚硝酸盐还原为一氧化氮

湖泊反硝化作用中的亚硝酸盐还原过程主要由变形菌和拟杆菌完成^[26-28]。亚硝酸盐还原为一氧化氮的反应可以由两种结构不同但功能相似的酶来实现:含血红素的 cd_1 亚硝酸盐还原酶(cd_1 -NIR)和含铜的亚硝酸盐还原酶(Cu-NIR)。这两种酶分别由 $nirS$ 和 $nirK$ 基因编码,其中 $nirS$ 基因的分布更广,存在于接近75%的反硝化微生物中^[22,29]。这两种酶都位于细胞周质,不能贡献能量保存^[23,28]。它们的区别在于催化亚硝酸盐还原产物的差异: cd_1 -NIR的反应产物往往是 NO ,而Cu-NIR在强还原和高pH条件下可能催化形成 N_2O ^[22]。

1.3 一氧化氮还原为氧化亚氮

一氧化氮作为反硝化的中间产物,具有细胞毒性,几乎不在胞内积累,很快在细胞内被还原或被微生物排泄到环境中^[22,30]。一氧化氮还原为氧化亚氮的过程是由一氧化氮还原酶(NOR)催化实现的,由 $norB$ 或 $norZ$ 基因编码^[31-33]。

N_2O 是一种强烈的温室气体,其释放速率与湖泊氮输入密切相关。McCrackin统计了26个不同类型(不同营养状态)湖泊的相关数据发现, N_2O 释放速率与湖泊硝态氮载荷呈显著正相关^[34]。本文在此基础上统计了全球180个不同营养状态湖泊的 N_2O 释放数据(涵盖了寡营养、中营养、富营养、重度富营养湖泊),发现 N_2O 释放速率有随湖泊富营养化逐渐升高的趋势(表2)。据估计,全球湖泊每年 N_2O 排放量为0.04–2.00 Tg N,约占内陆水体总排放量的12%–49%,到2050年 N_2O 排放量将增长至0.1–3.4 Tg N/yr^[34-35]。

湖泊 N_2O 释放速率具有很强的时空异质性,主要受到盐度、pH、DO、C/N比、 NO_3^- 浓度等环境因子的影响,具有季节性变化的特征。对反硝化微生物而言,非最适的生存条件(如高浓度氧气和低C/N比)往往会刺激其产生 N_2O ^[42]。盐度的升高和pH的降低也会刺激 N_2O 释放^[43-44]。高浓度DO(如5%)会抑制氧化亚氮还原酶的合成与活性从而增加 N_2O 的释放,但也不利于反硝化作用的进行,综合来看高浓度DO会导致反硝化过程中 N_2O/N_2 比增加^[45]。 N_2O 释放量往往随着C/N比升高而降低,可能与反硝化微生物的能量需求有关,在C/N比较高(即碳源充足)的情况下,异养的反硝化微生物倾向于实现 NO_3^- 到 N_2 的完全转化^[38,46]。 NO_3^- 作为反硝化作用的底物,其浓度与湖泊 N_2O 产量具有显著的正相关关系^[34]。 N_2O 释放表现出

表 2. 不同营养类型湖泊的 N₂O 释放速率
Table 2. N₂O emission rates in lakes of different nutrient types

Lakes	Descriptions	N ₂ O flux/[$\mu\text{mol N}/(\text{m}^2\cdot\text{d})$]			References
		System mean	Lake mean	Range	
121 small oligotrophic lakes	Oligotrophic		2.6	-7.6-13.8	[36]
Lake Makijarvi	Oligotrophic	2.3	2.3	-0.48-4.2	[37]
25 low-deposition region lakes	Oligotrophic		1.2	0-24	[34, 38]
Jankalaisenlampi pond	Mesotrophic		0.68		[37]
Kotsamolampi pond	Mesotrophic	14.8	-0.28		[37]
Lake Huahu	Mesotrophic		44	-38-191	[39]
24 high-deposition region lakes	Eutrophic		22.4	-24-163.2	[34, 38]
Lake Okaro	Eutrophic		1.2		[40]
Lake Postilampi	Eutrophic		4.9	-0.22-14	[37]
Lake Heinalampi	Eutrophic	18.8	-3.6		[37]
Lake Kevaton	Eutrophic		-2.16	-1.12--0.32	[37]
Lake Vehmasjarvi	Eutrophic		7	-0.26-11.8	[37]
Taihu Lake	Hyper-eutrophic	23.7	23.7	-3.4-109.8	[41]

Negative flux rates indicate movement of N₂O from the atmosphere to the lake.

明显的季节性特征实质上是受到温度和降水等因素的调控, 一般在夏季出现峰值^[47-48]。

1.4 氧化亚氮还原为氮气

微生物将氧化亚氮还原为氮气是减少氧化亚氮这种强大温室气体的主要途径。氧化亚氮还原酶(NOS, 由 *nosZDFYL* 基因编码)是目前已知的唯一能够催化此反应的酶, 广泛存在于变形菌门、拟杆菌门、绿菌门以及古菌中的泉古菌门和盐杆菌纲中^[49-50]。尽管能将硝酸根转化为氮气的完全反硝化菌更为常见, 对于某些缺少 NOS 的微生物来说氧化亚氮即是反硝化的最终产物(如 *Pseudomonas chlororaphis* 和某些 *Rhizobium*)^[22]。

氧化亚氮还原为氮气的速率主要受到两方面因素的影响, 即不完全反硝化菌的生态位分异(亚硝酸盐还原菌和氧化亚氮还原菌的分布不均衡导致局部氧化亚氮积累/不足)和完全反硝化菌的 NOS 酶活性(与其他反硝化酶相比, NOS 更加敏感, 其活性容易受到氧气、pH 以及硫化物的影响)^[50-53]。值得注意的是, 青藏高原盐湖中亚硝酸盐还原菌和

氧化亚氮还原菌群落组成的差异显著, 受到不同环境因子的调控, 其中氧化亚氮还原菌对环境变化更为敏感, 暗示着盐湖生境中可能是由不同微生物在实现完全反硝化过程(未发表数据)。这种不完全反硝化菌的差异是否会导致盐湖 N₂O 的局部积累或不足值得我们进一步探究。此外, 随着纬度的增加, 湖泊 N₂O 的排放速率逐级降低, 表现出从 N₂O 源转变为 N₂O 汇的趋势, 且作为 N₂O 汇的湖泊往往具有低 pH (<6.275)高 DOC (>7.492 mg/L)和低溶氧 (<7.805 mg/L)的特征^[54]。

1.5 湖泊反硝化速率及其影响因素

反硝化过程是湖泊氮去除的重要途径, 其去除的氮总量约占湖泊氮输入的 1%-36%^[15]。从全球范围来看, 湖泊的反硝化通量为 19-43 Tg/yr, 约占陆地氮输入的 7%-16%和全球反硝化通量的 4%-9%; 其中, 较小的湖泊(<50 km²)大约贡献了 2/3 的湖泊反硝化总量^[1]。

大部分反硝化作用发生在湖泊的下层滞水带及沉积物等低氧层位(一般为氧气浓度低于

0.2 mg/L 的次氧化层)。不同层位的反硝化速率差异很大,例如近岸表层沉积物反硝化速率为 22.5–366.2 nmol N/(L·h),湖泊水体的反硝化速率为 6.0–13.3 nmol N/(L·h)^[55–59]。沉积物反硝化过程的氮源一般由沉积物本身矿化产生的硝酸根提供,而不是上覆水体中的氮,因而大量无机氮会通过反硝化作用从沉积物中去除(76%–100%)^[15,56]。

从宏观上来看,湖泊生态系统反硝化速率的影响因素主要有水动力条件和氮载荷量^[1]。湖泊的水动力条件通过控制湖水中氮素的滞留时间来影响湖泊的反硝化速率。通常湖水滞留时间越长,参与反硝化过程的氮素比例越高,越有利于湖泊的反硝化作用。而湖泊氮载荷量决定了反硝化作用中底物的上限,通常湖泊的氮载荷量越高,其反硝化速率也越高^[60]。

从微观上来看,湖泊反硝化速率的影响因素主要有氧气浓度(影响酶活性)、底物(硝酸根和有机质)可利用性以及温度(季节性变化和昼夜变化)和盐度^[22]。本研究从文献中^[56,58–59]统计了长江中下游盆地 22 个富营养湖泊的反硝化速率、地化以及生物信息,使用 R 软件(Version 3.5.1) RandomForest (随机森林)程序包计算了不同环境因子对微生物反硝化速率的相对重要性。随机森林评价功能中的 Inc MSE 指数等价于 Mean Decrease Accuracy,描述的是对某个环境参数随机赋值时,随机森林预测准确度的降低程度,Inc MSE 越大表示环境因子的相对重要性越高(Inc MSE 等于 0 说明该环境参数与反硝化速率在统计学上不相关,Inc MSE 小于 0 则无统计学意义)^[61]。结果显示,对湖泊反硝化速率影响最大的前三个环境因子分别是沉积物硝酸盐浓度、溶解性总固体(可理解为盐度)和溶解氧(图 2),与前人的研究结论一致^[56, 58–59, 62–64]。

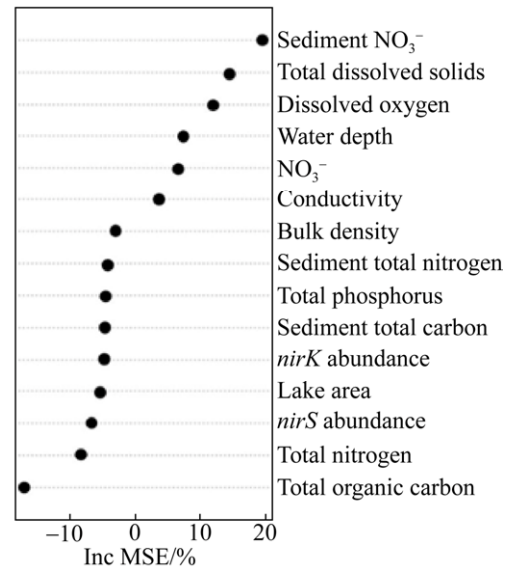


图 2. 环境因子对湖泊反硝化速率的相对贡献率
Figure 2. Relative contributions of environmental factors to lacustrine denitrification rates.

由于反硝化作用是一种高产能的异化还原过程,伴随着大量 ATP 的产生,因此能够耐受盐度胁迫,可在超盐环境(如盐湖,甚至盐度接近饱和的超盐湖)中发生^[63,65–66]。值得注意的是,高盐环境中反硝化微生物的生长似乎没有受到盐度的影响,但能够观察到反硝化终产物从氮气向氧化亚氮的转变^[65]。结合 NOS 比其他反硝化酶更加敏感的特征,该酶是否更容易受到盐度的抑制作用,使得盐湖有着更高的氧化亚氮释放速率,值得我们进一步探究。因此,全面评估盐湖的反硝化速率及其反硝化产物(氮气与氧化亚氮)比例对于全面评估全球氮循环具有重要意义。

2 与其他代谢过程耦合的反硝化新途径

2.1 共反硝化作用

早在 19 世纪末,学者们就发现微生物介导的

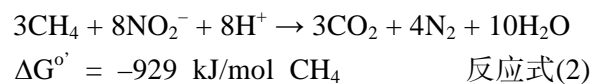
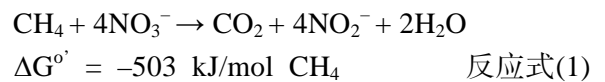
硝酸盐还原过程中会产生额外的氮气, 额外产生氮气的来源一直困扰着研究人员。直到 20 世纪末, 同位素标定技术的使用和共反硝化作用 (codenitrification) 的发现才回答了这个问题。共反硝化作用会产生杂交氧化亚氮和氮气^[18], 这种与传统反硝化作用共代谢的新途径通过微生物介导的亚硝化反应形成 N-N 键, 即亚硝酸盐或一氧化氮与其他含氮化合物(胺类、羟胺、铵根、叠氮化物、联氨等)各贡献一个氮原子生成氧化亚氮、氮气或同时生成两种气体(图 3)^[18, 67-69]。据现有研究, 共反硝化作用存在于细菌、古菌和真菌三域中, 但仅有为数不多的微生物能够执行该反应^[70-72]。共反硝化作用已被证实在土壤环境中扮演着重要的角色(如草地中 92% 的氮气由共反硝化作用产生^[73]), 而在水体环境中, 共反硝化过程的分布和速率仍有待进一步研究。前人研究显示, 湖水中氧化亚氮分子内氮同位素存在不一致的异常现象, 这暗示了水体环境中可能存在共反硝化作用^[74-75]。此外, 水体中浓度合适($\mu\text{mol/L}$ 级别)的氨基化合物也能够为共反硝化作用提供潜在底物^[76-77]。

基于共代谢的特征, 影响反硝化过程的环境因子对共反硝化作用也有着相似的影响, 包括氧气和有机碳底物的可利用性、pH 等^[68]。此外, 电子供体的类型很可能主导着共反硝化作用产物的种类(杂交氮气、杂交氧化亚氮或两种气体同时产生): 当电子供体为-3 价(如 R-NH_2 、 NH_3)时会生

成氮气; 当电子供体为-1 价(如 NH_2OH)时会生成氧化亚氮; 当电子供体为-2 价(如联氨)时会同时生成氧化亚氮和氮气^[18]。

2.2 厌氧甲烷氧化耦合反硝化

微生物介导的碳、氮元素循环过程通常耦合在一起, 并在很大程度上影响着湖泊的初级生产力、营养水平、温室气体释放等生态过程。近年来, 研究人员发现一种包含细菌和古菌的富集培养物能够将厌氧甲烷氧化过程与反硝化过程耦合在一起, 其中古菌 *Candidatus Methanoperedens nitroreducens* 属于厌氧甲烷氧化 ANME-2d 进化枝, 能够在完成厌氧甲烷氧化过程的同时, 将硝酸盐还原为亚硝酸盐; 细菌 *Candidatus Methyloirabilis oxyfera* 属于 NC10 门, 能够先将亚硝酸盐还原为 NO, 然后将 NO 歧化反应生成氮气和氧气, 再利用胞内的氧气氧化甲烷生成二氧化碳并释放能量, 该过程被称为亚硝酸盐依赖型厌氧甲烷氧化(Nitrite-dependent anaerobic methane-oxidation, N-DAMO), 古菌和细菌的反应式^[19, 78-80](反应式 1-2)。



以上两种反应统称为硝酸盐/亚硝酸盐依赖型厌氧甲烷氧化(nitrate/nitrite-dependent anaerobic

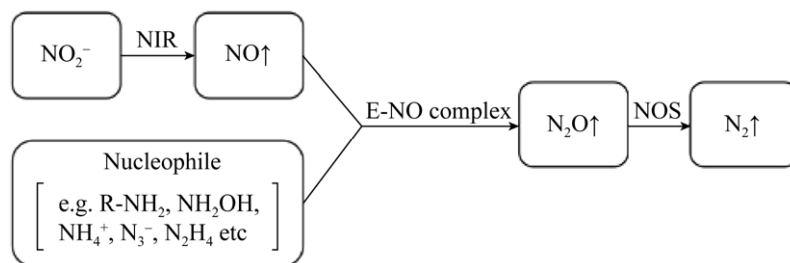


图 3. 共反硝化作用模式图(修改自参考文献[68])

Figure 3. Schematic diagram of codenitrification (modified after reference [68]).

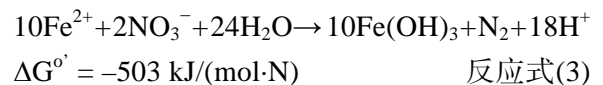
oxidization of methane, N-AOM)。N-AOM 微生物主要分布于富含甲烷和反应活性氮的低氧湖水或沉积物中, 尤其是好氧环境和厌氧环境的交界层位, 在这种环境中, 深层厌氧带扩散的氮容易被上覆水体中扩散的氧气氧化, 合适浓度的甲烷也能够充当电子供体促成反应的发生^[81]。同时, N-AOM 作为一种被忽视的甲烷汇, 缓解了温室气体甲烷的释放, 也促进了湖泊的氮去除, 具有十分重要的生态意义。Deutzmann 在针对 Lake Constance 湖氮循环的研究中发现, 该湖泊深层沉积物中 N-DAMO 过程消耗甲烷的速率为 1.8–3.6 nmol CO₂/(mL·d), 对沉积物总甲烷氧化过程的贡献能够达到约 5%^[82]。湖泊及其他水生环境中 N-DAMO 的速率和贡献见表 3。

影响 N-AOM 过程的环境因素包括温度、氧气、硝酸根、亚硝酸根、甲烷、有机碳、盐度等。N-AOM 微生物生长缓慢, 倍增时间约为 1–2 周, 因此更加倾向存在于湖泊中相对稳定的层位^[86]。大量研究表明, 硝酸根、亚硝酸根、甲烷等底物浓度及其配比控制着 N-AOM 微生物在湖泊生境中的分布和活力, 例如有机碳浓度能够通过控制产甲烷菌的活力对 N-AOM 过程产生实质性的影响^[83–84,87–89]。盐度则是 N-AOM 微生物群落结构和速率的重要限制因素, 前人统计了全球不同生境中 N-AOM 细菌的分布, 发现盐度是 N-AOM 细菌

群落组成的主控因素之一^[90]。值得注意的是, 前人研究显示 N-AOM 细菌能够存在于河口、盐湖等高盐生境(盐度高达 84 g/L)中并完成反应, 这暗示着嗜盐 N-AOM 细菌的存在^[88,91–92]; 而 N-AOM 古菌还没有在盐度超过 20 g/L 的环境中报道过^[93]。因此能够耐受更高盐度的 N-AOM 古菌是否存在仍有待进一步研究。

2.3 硝酸盐依赖型铁氧化

硝酸盐依赖型铁氧化(anaerobic nitrate-dependent Fe(II) oxidation, ANDFO)主要发生在无光照、缺氧的中性沉积环境中, 是一种微生物介导的、硝酸盐还原耦合溶解态或非溶解态铁氧化的反应, 在铁、氮循环中起着不可忽视的作用(反应式 3)^[20, 94–96]。



前人发现多种不同营养型的细菌和古菌都能够执行 ANDFO 过程, 它们大部分属于异养或混合营养型的变形菌(如 *Acidovorax* sp. Strain 2AN、*Dechloromonas* sp. Strain UWNR4), 只有极少数细菌(如 *Pseudogulbenkiania* sp. strain 2002、富集培养物 KS)和一株嗜热古菌(*Ferroglobus placidus*)属于无机自养型微生物^[94,97–100]。

湖泊中微生物介导的硝酸盐依赖型铁氧化过程具有重要的生态功能。首先, 硝酸盐还原和

表 3. 水生生态系统 N-DAMO 速率范围及甲烷氧化贡献率

Table 3. N-DAMO rate ranges and their methane oxidation contribution to aquatic ecosystems

Ecosystems	Descriptions	Rates	Contributions/%	References
Lake constance	Oligotrophic	1.8–3.6 nmol CO ₂ /(mL·d)	5	[82]
Freshwater pond	Nitrate levels of 1–2 mmol/L	2.3 mmol/(m ² ·d)	47	[83]
Xiazuhuhu wetland	Eutrophic	0.2–14.5 nmol CO ₂ /g dry soil·d	2.7–4.3	[84]
Mangrove wetland	Salinity ~8g/L	25.93–704.80 nmol CO ₂ /g dry soil·d	\	[85]

“\” means no reference data.

铁氧化的耦合过程形成的三价铁矿物,能够改变营养物质或重金属的流动性进而影响局部环境^[101-102]。其次,铁浓度可以直接影响多种氮循环步骤,例如生物可利用二价铁浓度的增加能加快河口沉积物中硝酸盐向氨的转化速率[由 14.9 升高至 25.5 $\mu\text{mol Fe}/(\text{L}\cdot\text{d})$, 盐度为 26.8 g/L],从而增加 10%–20%的氮去除率,并降低反硝化速率^[103]。此外,Hauck 的研究显示, Lake Constance 湖沉积物环境中的 ANDFO 微生物占总硝酸盐还原微生物的 58%^[104]。这也从另一方面说明了 ANDFO 过程是湖泊氮循环中不容忽视的一部分。

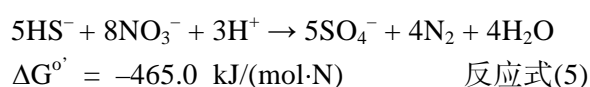
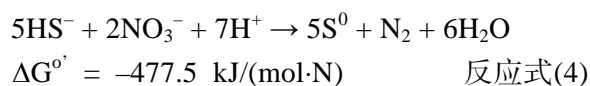
ANDFO 过程受到硝酸根浓度、有机底物、磷酸根浓度等环境因素的影响。前人的实验表明,大多数混合营养型的 ANDFO 微生物需要一定浓度的有机底物来维持最佳的铁氧化速率^[20,95,105-106];磷酸根会与亚铁形成磷酸亚铁沉淀(蓝铁矿),显著降低亚铁的生物可利用性,从而大幅降低 ANDFO 速率;而碳酸根在过饱和情况下虽然不会形成沉淀,也几乎不改变 ANDFO 速率,但会影响 ANDFO 过程最终形成的矿物种类^[107]。

ANDFO 过程会释放大量的自由能(503 kJ/(mol·N)),能帮助微生物抵御高渗透压胁迫^[103]。Emmerich 在一个超盐湖(盐度为 348.6 g/L)沉积物中检测到 210 cells/g(沉积物干重)硝酸盐还原铁氧化菌,说明 ANDFO 过程也许能够在接近盐度饱和的条件下发生^[108]。以上发现均暗示着 ANDFO 过程可能具有一定的耐盐性,但未知其如何响应盐度变化。

2.4 硝酸盐还原硫氧化

除了有机碳和二价铁,硝酸盐还原菌还能够以硫化物为电子供体完成硝酸盐还原过程。在该过程中,硝酸根通过反硝化途径还原为氮气,硫

化物可以被氧化为硫单质储存于细胞内或者进一步氧化为硫酸根。根据产物的不同,反应表达式不同(反应式 4–5)。



硝酸盐还原硫氧化微生物主要为 α 、 β 、 γ 和 ϵ 变形菌纲。作为最常见的硝酸盐还原硫氧化菌, *Thiobacillus denitrificans* 和 *Sulfurimonas denitrificans* 广泛分布于湖泊、海洋等自然环境^[21,109]。一些 γ 变形菌纲的硝酸盐还原硫氧化菌(如 *Thioalkalivibrio*、*Thiohalomonas*、*Thiohalophilus*、*Thioalkalispira*、*Thiohalorhabdus*) 在盐度高达 24% 的盐湖环境中被分离出来,证明这些菌具有很高的耐盐能力^[110-113]。从热力学角度看,伴随着质子释放的硫氧化是一种高产能反应(每传递一个电子产生 84.7–88.4 kJ/mol 能量),因此该反应能够产生足够的能量抵御盐度压力,这与纯菌分离试验的结果相吻合^[66,114]。

硫化物丰富的海洋是最早发现硝酸盐还原硫氧化途径的生境。但前人研究显示,在淡水生态系统中该途径也扮演着重要的角色^[21,115-116]。硫化氢和硝酸根在水体中富集都会引起严重的环境问题,微生物介导的硝酸盐还原硫氧化反应能够同时去除水体中的氧化态氮和还原态硫,在维持水环境稳定方面具有重要的生态意义。例如,在西班牙三个淡水湖中,硝酸盐还原硫氧化途径对湖泊氮去除过程的贡献可达 25%–40%;珠江河口沉积物环境中,硝酸盐还原硫氧化速率为 0.196–0.903 mmol/(L·h),在硝酸盐较低时(<0.2 mmol/L),硝酸盐还原硫氧化途径对氮去除过程的贡献率则能够达到 26%^[21,117]。

3 总结和展望

综合湖泊反硝化微生物群落与反硝化速率的研究发现,微生物介导的湖泊反硝化过程在减轻湖泊氮载荷方面起着不可忽视的作用,具有重要的生态环境意义。此外,微生物介导的反硝化与其他元素循环的耦合对于湖泊氮去除也有实质性贡献,拓宽了湖泊氮循环的研究思路。宏基因组、宏转录组、稳定同位素标记等新兴技术手段的出现与应用,帮助学者们更清晰、准确地认识了反硝化微生物代谢机理和湖泊反硝化通量,也发现了一些亟待解决、值得深入探究的科学问题,例如:(1)人类活动引起全球氮沉降增加如何影响湖泊尤其是脆弱的高原湖泊反硝化微生物群落结构和功能;(2)目前的研究大多报道了某一种反硝化途径的脱氮速率与贡献率,缺乏对同一湖泊各种反硝化途径脱氮贡献的整体认识;(3)反硝化微生物与环境相互作用的生理机制尚不明确,有待更深入的研究;(4)氧化亚氮还原酶 NOS 比其他反硝化酶更敏感,那么在盐湖、碱湖中能否观察到 N_2O 在反硝化终产物中占比升高,仍有待研究;(5)反硝化过程与其他元素循环耦合的酶学机制尚不明确,这一机制的突破将有助于学者们利用基因手段发现和定量耦合反应的微生物群落,加快湖泊元素循环的研究进程。湖泊生境在全球氮循环中扮演着重要角色,湖泊生态功能与物质循环速率一直是国内外研究的热点话题。然而目前人类活动的加剧使得湖泊生态功能表现出退化的趋势,因此深入探究湖泊的物质循环规律能够为恢复湖泊生态功能、维持其生态稳定提供重要的参考意见。

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Progress in microbially mediated denitrifying in lacustrine environments

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Abstract: Microbially mediated denitrification in lacustrine ecosystems has profound impact on regional and even global environment and climate change. Thus, studying lake microbial denitrification process and rate helps us obtain comprehensive understanding of nitrogen biogeochemical cycle in lacustrine ecosystems and its role in the global nitrogen cycle. This review summarizes the denitrification processes, rates and involved microbial community compositions in lacustrine ecosystems and their influencing factors. The covered denitrification processes include typical denitrification processes and those coupled with other elemental cycles including co-denitrification coupled with organic nitrogen, nitrate/nitrite-dependent anaerobic oxidization of methane coupled with carbon cycle, anaerobic nitrate-dependent Fe(II) oxidation coupled with iron cycle, nitrate/nitrite-dependent sulfide oxidation coupled with sulfur cycle. Finally, current researches and future directions related to denitrification processes in lacustrine environments were summarized.

Keywords: lakes, microbes, denitrification, rate, influencing factors

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